

## Further Considerations of the Provenance of the Rocourt Tephra: Volcanic Mafic Minerals and Age

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### Abstract

The Rocourt Tephra (RT) is a pyroclastic fallout deposit that was discovered in Belgium three-quarters of a century ago. Since then, the search for its source volcano has been in vain. Recently, however, two volcanoes of the West Eifel Volcanic Field have been put forward, namely the Dreiser Weiher and the Pulvermaar, but it has been shown by the geochemical fingerprints of associated pyroxenes that neither of them is compatible with the RT (Juvigné et al., 2024, DOI: 10.1007/s00445-024-01756-2). We present here additional arguments that support this conclusion. They arise from the comparison of volcanic mafic mineral associations of the RT and the Dreiser Weiher and Pulvermaar tephras. Furthermore, details of the stratigraphic study that led to the age estimate of between 70 and 80 ka for RT are reported, in order to widen the scope for determining the ages of volcanoes or tephras in the Eifel Volcanic Field (or elsewhere).

**Keywords:** Rocourt Tephra, volcanoes, Dreiser Weiher, Pulvermaar, Eifel Volcanic Field – Germany, Eifel – Belgium, stratigraphy, Quaternary, Upper Pleistocene

### Résumé

**Considérations complémentaires relatives à la provenance du Téphra de Rocourt : minéraux mafiques volcaniques et âge.** Le Téphra de Rocourt est une retombée de cendre volcanique qui a été découverte en Belgique, il y a trois quarts de siècle. Depuis lors, toutes les recherches destinées à découvrir son volcan émetteur ont été vaines. Deux volcans de l'Eifel occidentale ont été cités, à savoir le Dreiser Weiher et le Pulvermaar, mais il a été démontré par la composition chimique des pyroxènes qu'ils sont incompatibles avec le Téphra de Rocourt

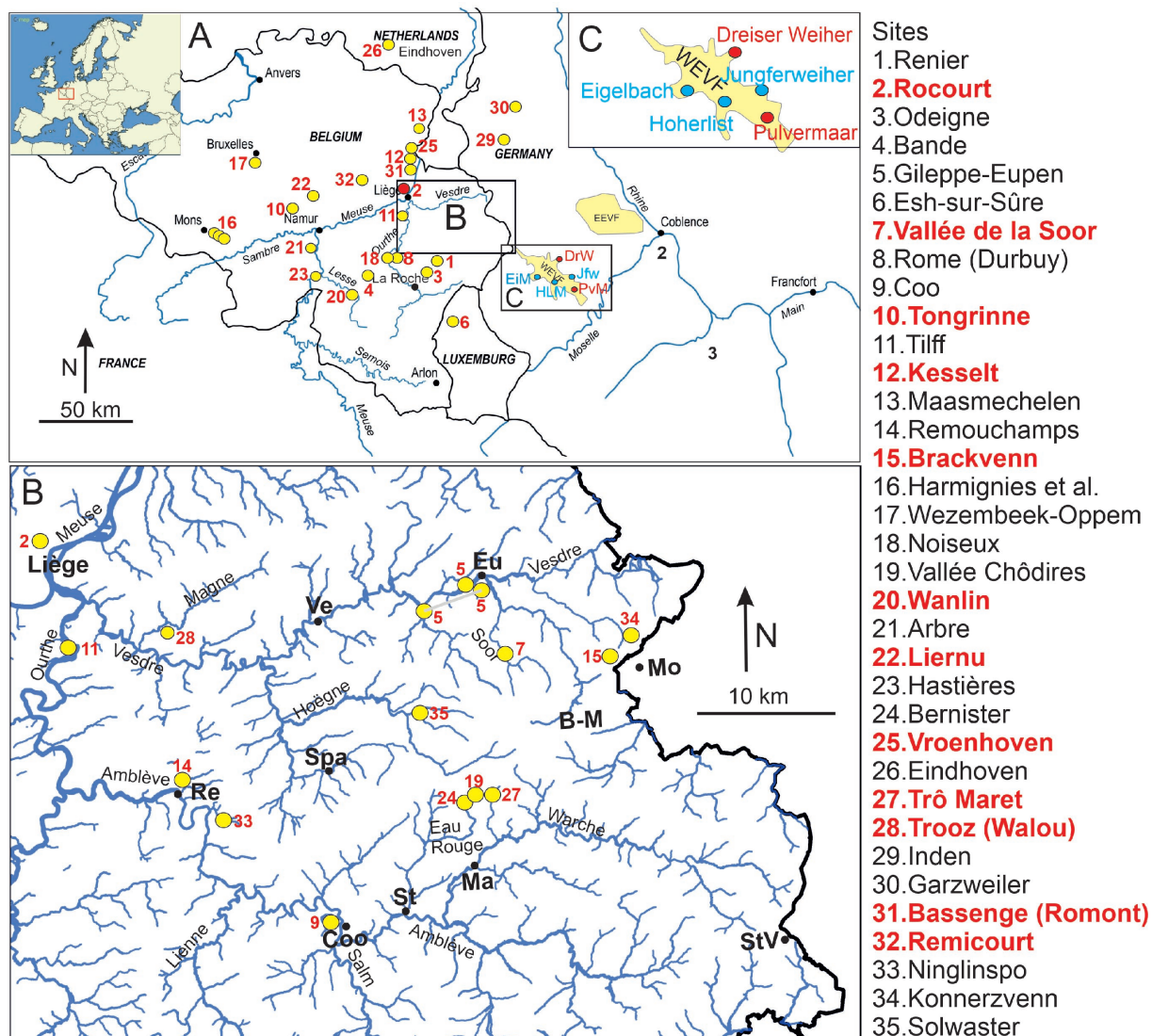
(Juvigné et al., 2024, DOI: 10.1007/s00445-024-01756-2). Le présent travail présente des arguments supplémentaires qui supportent cette conclusion ; ils relèvent de la comparaison des associations de minéraux mafiques volcaniques des trois parties impliquées. Par ailleurs, les détails de l'étude stratigraphique qui ont conduit à estimer l'âge du Téphra de Rocourt entre 70 et 80 ka sont rapportés, de façon à élargir la tolérance avec d'éventuelles datations de volcans de l'Eifel ou d'ailleurs.

**Mots-clés :** Téphra de Rocourt, volcans, Dreiser Weiher, Pulvermaar – Allemagne, Eifel – Belgique, stratigraphie, Quaternaire, Pléistocène supérieur

## 1. Introduction

The enstatite bearing volcanic ash deposit (now known as Rocourt Tephra, RT, see below) was discovered in a dispersed state (i.e., as a cryptotephra) in reworked deposits of Upper Belgium (Gullentops, 1952; Tavernier and Laruelle, 1952), then in a paleosol of the Eemian–Weichselian transition of the Rocourt loess stratotype (Gullentops, 1954). Since then, it has been found in about 35 sites in Belgium, the Grand-Duchy of Luxembourg, the Netherlands and the Lower Rhine Bay/Germany (Fig. 1; Table 1). In all cases, its identification is based on the volcanic mafic mineral (vmm) assemblage as determined using a polarizing microscope, and more particularly by a high frequency of enstatite and brown amphibole. It became the Tuf de Rocourt (Juvigné, 1977b) then the Rocourt Tephra (Juvigné, 1991). The optical determination of the enstatite was confirmed by microprobe analyses by Bustamante-Santa Cruz (1973) and that of the clinopyroxenes and the amphibole by Juvigné (1990).

When the volcanic ash was first discovered by Gullentops (1952), its origin from the West Eifel Volcanic Field (WEVF) was suspected by the author, but, to date, the emitting volcano has never been identified. Following the identification of the RT in a loess section of the Lower Rhine Bay, Gullentops and von der Hocht (1998) pointed out that the Dreiser Weiher volcano was a possible source on the sole basis of its large size and its relative proximity in the northern part of the WEVF. This hypothesis has not received any further attention. Lenaz et al. (2010) identified a tephra containing traces of enstatite in the Jungferweiher lake core (WEVF), and they clearly correlated it with the RT. Förster et al. (2020) took up the above correlation by integrating a tephra also containing traces of enstatite in two other lacustrine sequences of the WEVF, namely Eigelbach and Hoherlist maare. By designating the Pulvermaar as the source volcano for the RT occurrences in these three maare (Jungferweiher, Eigelbach, Hoherlist), the authors provoked a comparative study of the RT with the proximal tephra of Pulvermaar (Juvigné et al., 2024) This study also included the Dreiser Weiher. Juvigné et al. (2024) demonstrated by the study of the chemical composition of the pyroxenes that neither the Pulvermaar nor the Dreiser Weiher could be the source of the RT. In their paper, the optical determinations of the vmm had little weight because of the variations inherent in the complexity of the factors that can modify, at least quantitatively, the mineralogical association of the same tephra. Those factors can include the heterogeneity of the magma chamber, the alteration in the host terrains, as well as air transport and laboratory practices. Nevertheless, the results obtained by Juvigné et al. (2024) support the conclusion obtained from the chemical composition of the pyroxenes.



**Figure 1:** Location of sites where the volcanic mafic minerals (vmm) of the Rocourt Tephra (RT) have been identified as well as some localities where enstatite has been recognized as traces in various associations. Explanation: the numbers referring to localities correspond to the relevant literature references listed in Table 1; sites in bold and red font are of major interest, providing mineralogical and/or chemical and/or stratigraphical data. Abbreviations: WEVF = West Eifel Volcanic Field; PvM = Pulvermaar; DrW = Dreiser Weiher; Jfw = Jungferweiher; HIM = Hoherlist Maar; EiM = Eigelbach Maar; EEVF = East Eifel Volcanic Field; B-M = Baraque Michel; Eu = Eupen; Lg = Liège; Ma = Malmedy; Mo = Monjoie; Re = Remouchamps; St = Stavelot; StV = Saint-Vith; Ve = Verviers.

**Table 1:** For the Rocourt Tephra, articles reporting the presence of the RT in deposits at the sites shown in Fig. 1.

| <b>Authors</b>         | <b>Year</b> | <b>Site #</b> | <b>Host sediment</b>                |
|------------------------|-------------|---------------|-------------------------------------|
| Tavernier and Laruelle | 1952        | n/a           | Alluvial plain                      |
| Gullentops             | 1952        | 1             | Soil                                |
| Gullentops             | 1954        | 2             | Loess section                       |
|                        |             | 3 & 4         | Slope deposits                      |
|                        |             | 5             | Slope deposits & terrace            |
| Bourguignon            | 1955        | 1             | Soil                                |
| Hermans                | 1955        | 6             | Slope deposits                      |
| Bastin et al.          | 1972        | 7             | Valley bottom: periglacial deposits |
| Juvigné and Mullenders | 1972        | 8             | Loess section                       |
| Bustamante Santa Cruz  | 1973, 1974  | 9             | Terrace                             |
| Juvigné                | 1973        | 10            | Loess section                       |
|                        |             | 11            | Terrace                             |
| Juvigné                | 1974        | 12            | Loes section                        |
| Pissart                | 1974        | 13            | Terrace                             |
| Ek                     | 1974        | 14            | Cave deposits                       |
| Bastin et al.          | 1974        | 15            | Rampart of lithalsa                 |
| Pissart et al.         | 1975        | 7             | Valley bottom: periglacial deposits |
| Juvigné                | 1976        | 2             | Loess section                       |
|                        |             | 12            | Loess section                       |
|                        |             | 10            | Loess section                       |
|                        |             | 16            | Loess section                       |
|                        |             | 16            | Loess section                       |
|                        |             | 16            | Short term excavation               |
|                        |             | 17            | Short term excavation               |
| Juvigné                | 1977b       | n/a           | Varia                               |
| Juvigné                | 1977a       | 18            | Loess & eolian sand                 |
| Bastin and Juvigné     | 1978        | 19            | Valley bottom: periglacial deposits |
| Juvigné                | 1979a       | 20            | Loess section                       |
| Juvigné                | 1979b       | s.o.          | Terrace                             |
| Juvigné and Pissart    | 1979        | 15            | Slope deposits                      |
| Quinif et al.          | 1979        | 21            | Cave deposits                       |
| Bolline et al.         | 1980        | 22            | Loess section                       |

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| <b>Authors</b>               | <b>Year</b> | <b>Site #</b> | <b>Host sediment</b>                |
|------------------------------|-------------|---------------|-------------------------------------|
| Pissart and Juvigné          | 1980        | 15            | Rampart of lithalsa                 |
| Ballmann et al.              | 1980        | 23            | Cave deposits                       |
| Haesaerts et al.             | 1981        | n/a           | Loess section                       |
| Pissart and Juvigné          | 1982        | 24            | Valley bottom: periglacial deposits |
| Meijs and de Lang            | 1983        | 25            | Loess section                       |
|                              |             | 26            | Loess section                       |
| Mees and Meijs               | 1984        | 25            | Loess section                       |
| Juvigné and Mörner           | 1984        | n/a           | Palaeolake                          |
| Juvigné                      | 1985        | 27            | Terrace                             |
| Juvigné                      | 1990        | n/a           | Varia                               |
| Pouclet and Juvigné          | 1993        | 27            | Loess section                       |
| Lacroix                      | 1993        | 28            | Cave deposits                       |
| Juvigné                      | 1993        | n/a           | Varia                               |
| Juvigné et al.               | 1996        | 12            | Loess section                       |
| Gullentops and von der Hocht | 1998        | 29 & 30       | Loess section                       |
| Meijs and Groenendijk        | 1999        | 12            | Loess section                       |
| Bringmans et al.             | 1999-2000   | n/a           | Loess section                       |
| Renson et al.                | 2002        | 28            | Cave deposits                       |
| Meijs                        | 2002        | 12 & 25       | Loess section                       |
| Pirson et al.                | 2004        | 28            | Cave deposits                       |
| Juvigné et al.               | 2008        | 31            | Loess section                       |
| Pouclet et al.               | 2008        | 2, 12, 32     | Loess section                       |
|                              |             | 28            | Cave deposits                       |
|                              |             | 27            | Terrace                             |
| Rixhon and Juvigné           | 2010        | 33            | Valley bottom: periglacial deposits |
| Meijs                        | 2011        | 10            | Loess section                       |
| Pirson and Juvigné           | 2011        | 28            | Cave deposits                       |
| Juvigné et al.               | 2013        | 32            | Loess section                       |
| Juvigné                      | 2016        | 34            | Rampart of lithalsa                 |
| Jouannic et al.              | 2016        | 2             | Loess section                       |
| Haesaerts et al.             | 2016        | 16            | Loess section                       |
| Juvigné et al.               | 2022        | 35            | Valley bottom: periglacial deposits |

Furthermore, in tephra correlation research, the age of a deposit constitutes a valued guide, and it has been used by Lenaz et al. (2010) and Förster et al. (2020) for the RT. However, the RT has never been found in a primary position. Despite the progress of loess stratigraphy in Belgium (Haesaerts et al., 2016), the vertical distribution of its vmm in the Lower Weichselian loess sequences weakens the narrowness of the estimate currently proposed between 78 and 80 ka. Detailed data on this subject will be provided below to illustrate this reservation.

## **2. Analytical Procedure**

Dense minerals ( $\delta > 2.8$ ) were separated as follows: boiling in  $\text{HCl}_{10\% \text{vol}}$ ; sieving by 355/75  $\mu\text{m}$ ; extraction of dense minerals with bromoform ( $\delta = 2.8$ ) in a separating funnel by repeating agitation–decantation–harvest cycles, until no more harvest was obtained (generally three to five cycles). Aliquots were examined under the binocular magnifier, and smear slides were prepared in Canada balsam for identification with a petrological microscope.

## **3. Volcanogenic Mafic Minerals of the Rocourt Tephra**

The association consists of orthopyroxenes, clinopyroxenes, brown amphiboles, titaniferous magnetite and Cr-spinel. Due to several factors including the zonation of magma and the fallout (natural factors), as well as from sampling to optical determination (technical factors), the frequency ranges of the individual minerals are somewhat large. In the sites with high frequencies of vmm, a characteristic mineral suite was calculated (Table 2). Photographs of the three most common volcanogenic mafic minerals of the Rocourt Tephra are presented in Figure 2.

## **4. The Dreiser Weiher Tephra**

The location of the volcano and the detailed positions of the samples stratigraphically are available in Juvigné et al. (2024, Fig. 3). Two populations share some 90% of the entire population: very dark green to black euhedral minerals coated with dark magmatic glass and colorless to greenish grains which are not coated with magmatic glass. The magmatic coating impairs the transparency during microscope examination. Nevertheless, numerous dark green euhedral minerals could be identified as clinopyroxene as well as agglomerates of small euhedral clinopyroxenes. The colorless crystals consist of olivine. Traces of another two minerals were found: anhedral lawn-green crystals with serrated edges (clinopyroxene) and prismatic brown minerals (amphiboles). No significant stratigraphic variation was observed in the sequence (Table 3). Photographs of the three most common volcanogenic mafic minerals of the Dreiser Weiher Tephra are presented in Figure 3.

## **5. The Pulvermaar Tephra**

The location of the volcano and the detailed position of the samples are given in Juvigné et al. (2024, Fig. 4).

**Table 2:** Percentage ranges of the vmm of the Rocourt Tephra. *Explanation:* Cpx = clinopyroxene; Ens = enstatite; Amp = amphibole; *n* = the number of grains counted; n.a. = not available.

| Locality                             | Determinator     | Ref. | Cpx  | Ens  | Amp  | Spinel | <i>n</i> |
|--------------------------------------|------------------|------|------|------|------|--------|----------|
| Eupen, Soor valley, one section      | Juvigné          | [1]  | 53   | 27   | 20   |        | n.a.     |
| Eupen, Soor valley, various sections | Juvigné          | [2]  | 50   | 25   | 25   |        | n.a.     |
| Rocourt, loess section               | Juvigné          | [3]  | 58.7 | 9.4  | 31.8 |        | 388      |
| Kesselt, loess section               | Juvigné          | [3]  | 65.5 | 13.4 | 21.1 |        | 739      |
| Tongrinne, loess section             | Juvigné          | [3]  | 49.7 | 25.6 | 24.6 |        | 107      |
| Wanlin, loess section                | Juvigné          | [4]  | 28.9 | 43.1 | 28   |        | 2888     |
| Liernu, loess section                | Juvigné          | [5]  | 22.9 | 34.9 | 42.9 |        | 73       |
| Hautes Fagnes, Trô Maret, terrace    | Juvigné          | [5]  | 46.8 | 22   | 31.2 |        | n.a.     |
| Vroenhoven, loess section            | Meijs            | [6]  | 76.7 | 8.7  | 13.8 | 0.8    | n.a.     |
| Hautes Fagnes, Trô Maret, terrace    | Juvigné          | [7]  | 45.3 | 25.5 | 29.1 |        | 1300     |
| Trooz, Walou, cave                   | Pirson & Juvigné | [8]  | 45.4 | 40.1 | 14.3 |        | 1279     |
| Bassenge, Romont, loess section      | Juvigné          | [9]  | 90.9 | 3.8  | 5.2  |        | 10550    |
| Remicourt, loess section             | Juvigné          | [10] | 67.2 | 10   | 22.7 |        | 3735     |
| Min.                                 |                  |      | 22.9 | 3.8  | 5.2  |        |          |
| Max.                                 |                  |      | 90.9 | 43.1 | 42.9 |        |          |

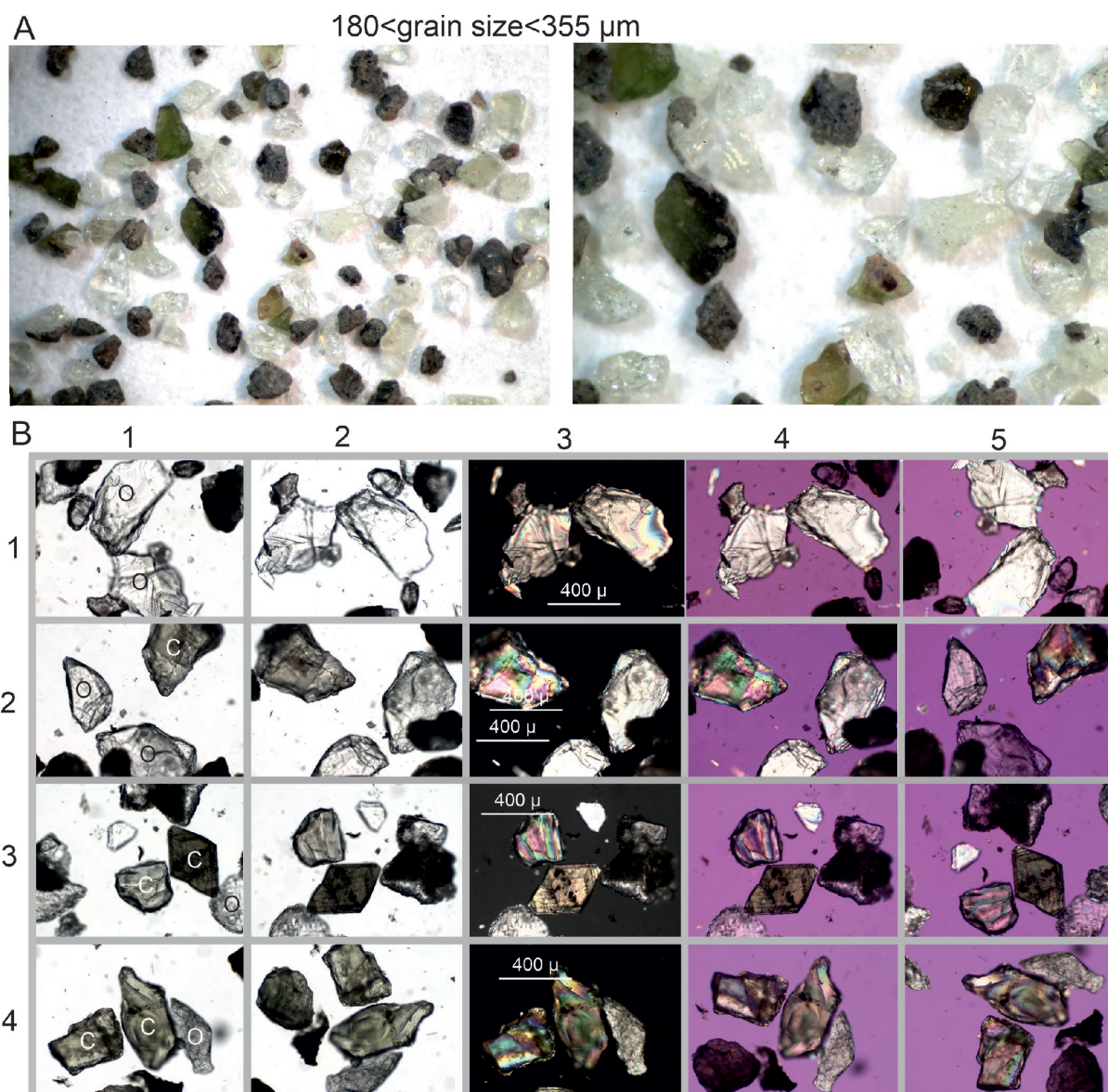
References: [1] – Bastin et al. (1972); [2] – Pissart et al. (1975); [3] – Juvigné (1976); [4] – Juvigné (1979a); [5] – Bolline et al. (1980); [6] – Meijs and de Lang (1983); [7] – Juvigné (1985); [8] – Pirson and Juvigné (2011); [9] – Juvigné et al. (2008); [10] – Juvigné et al. (2013).

**Table 3:** Frequency (%) of the volcanogenic mafic mineral suite of the Dreiser Weiher Tephra after optical determinations (magnifier and microscope). *Explanation:* *n* = number of mineral grains counted.

| Label            | Clinopyroxene | Enstatite + amphibole | Olivine | <i>n</i> |
|------------------|---------------|-----------------------|---------|----------|
| Dreiser Weiher 2 | 38.5          | 0                     | 61.5    | 117      |
| Dreiser Weiher 3 | 40.0          | 2.9                   | 57.1    | 105      |
| Dreiser Weiher 4 | 46.7          | 0                     | 53.3    | 112      |



**Figure 2:** Photographs of the three most common volcanogenic mafic minerals of the Rocourt Tephra. *Rows:* E = enstatite; A = amphibole; C = clinopyroxene. *Columns:* (1) and (2) orthogonal positions in plane-polarized light to highlight pleochroism; (3), (4), (5) maximum illumination between cross-polarized light to highlight birefringence with intercalated gypsum (4) and (5).



**Figure 3:** Photographs of the two most common volcanogenic mafic minerals of the Dreiser Weiher Tephra. C = cpx; O = olivine.

(A) Under magnifier: dark grains are clinopyroxenes; colorless to pale greenish grains are olivines.

(B) Rows: C = clinopyroxene; O = olivine. Columns: (1) and (2) orthogonal positions in plane-polarized light to highlight pleochroism; (3), (4), (5) maximum illumination between cross-polarized light to highlight birefringence with intercalated gypsum in (4) and (5).

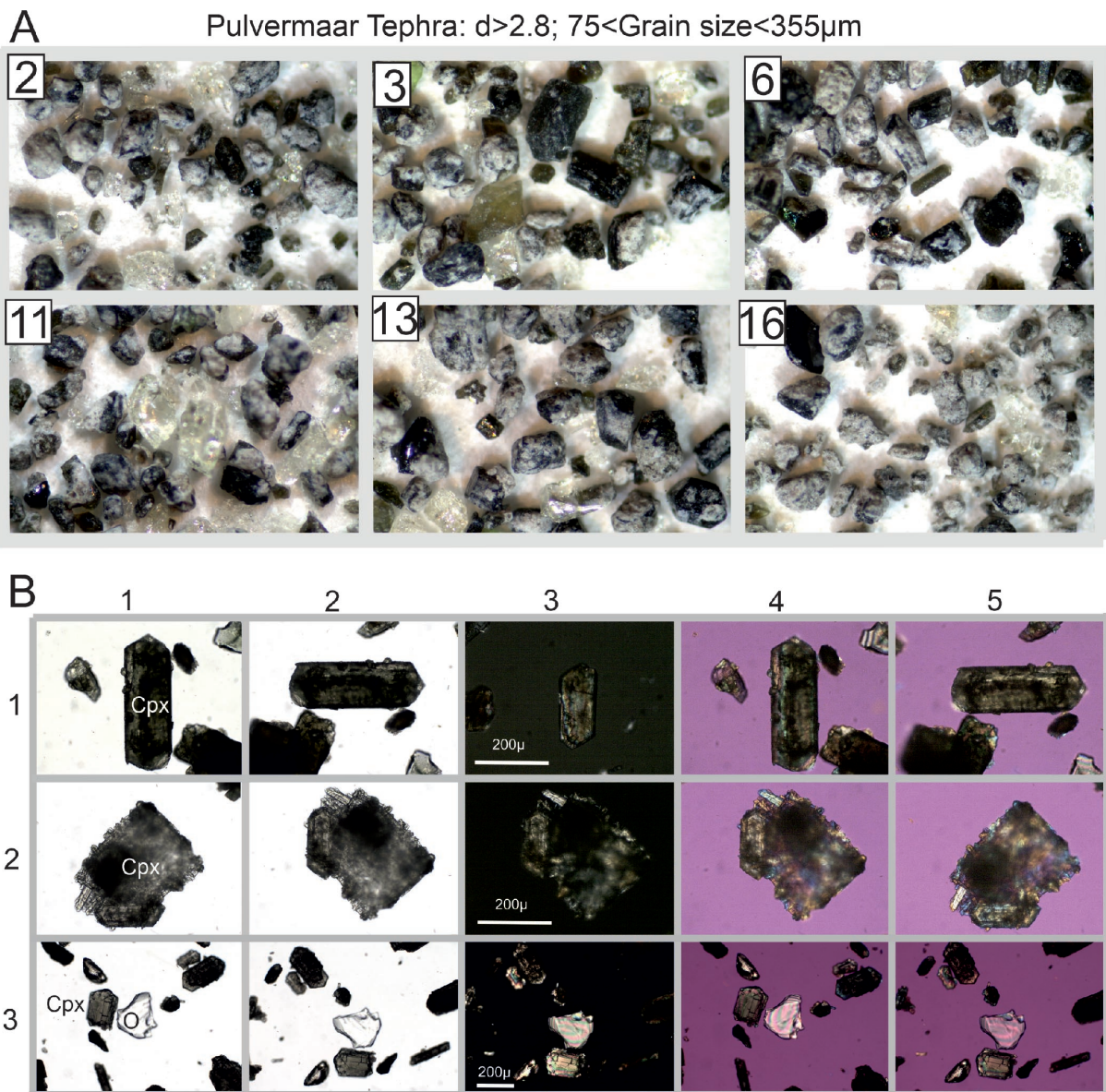
**Table 4:** Frequency (%) of the volcanogenic mafic mineral suite of the Pulvermaar Tephra after optical determinations (magnifier and microscope). *Explanation:*  $n$  = number of mineral grains counted.

| Label         | Clinopyroxene | Enstatite +<br>amphibole | Olivine | $n$ |
|---------------|---------------|--------------------------|---------|-----|
| Pulvermaar 1  | 88.7          | 0                        | 11.3    | 106 |
| Pulvermaar 2  | 76.1          | 2.2                      | 21.7    | 105 |
| Pulvermaar 3  | 92.7          | 0                        | 7.3     | 123 |
| Pulvermaar 4  | 95.3          | 0                        | 4.7     | 128 |
| Pulvermaar 5  | 96.2          | 1.9                      | 1.9     | 106 |
| Pulvermaar 6  | 97.1          | 0                        | 2.9     | 134 |
| Pulvermaar 7  | 98.3          | 0                        | 1.7     | 116 |
| Pulvermaar 8  | 93.4          | 1.6                      | 4.9     | 122 |
| Pulvermaar 9  | 91.8          | 2                        | 6.1     | 105 |
| Pulvermaar 10 | 88.4          | 11.6                     | 0       | 109 |
| Pulvermaar 11 | 81.8          | 18.2                     | 0       | 110 |
| Pulvermaar 12 | 94.8          | 5.2                      | 0       | 144 |
| Pulvermaar 13 | 88.6          | 8.5                      | 2.9     | 105 |
| Pulvermaar 14 | 87.5          | 12.5                     | 0       | 112 |
| Pulvermaar 15 | 96.8          | 13.2                     | 0       | 126 |
| Pulvermaar 16 | 94.9          | 5.1                      | 0       | 118 |
| Pulvermaar 17 | 89.6          | 10.4                     | 0       | 134 |

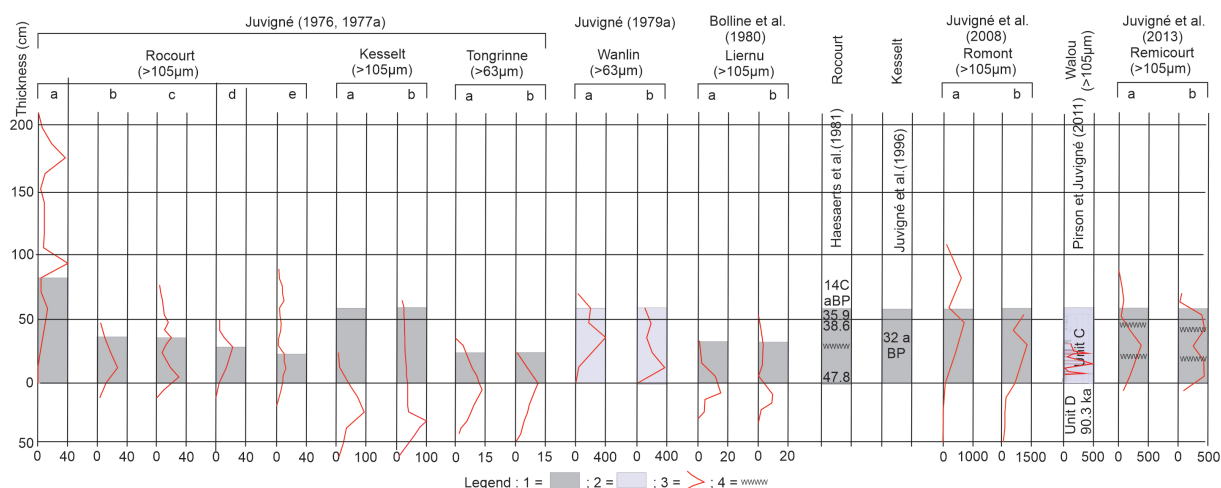
Two populations make up most of the mass: very dark green to black grains mainly coated with magmatic glass are largely dominant over transparent colorless grains. Under the microscope, magmatic glass coating impairs the transparency of the former minerals. Nevertheless, several of them could be identified as subhedral to euhedral clinopyroxene as well as agglomerates of small euhedral clinopyroxenes. The transparent colorless minerals are likely to be olivine. Otherwise, a few anhedral lawn-green minerals with serrated edges are present (clinopyroxene) as well as traces of prismatic brown grains (amphibole). The latter two minerals and the olivines are not coated with magmatic glass and so are grains of coarser crystals. Qualitative examination of the eighteen samples has been done and no significant stratigraphical variation was observed (Table 4). Photographs of the two most common volcanogenic mafic minerals of the Pulvermaar Tephra are presented in Fig. 4

## 6. Stratigraphical Distribution of the Rocourt Tephra

In sites where the Upper Pleistocene lithostratigraphic units are sufficiently developed, the vertical distribution of the vmm was investigated, with the aim of finding the RT layer visible to the naked eye at a peak of concentration. However, the objective of identifying the RT in its primary position has never been achieved. In some cases, secondary concentration peaks have



**Figure 4:** Photographs of the two most common volcanogenic mafic minerals of the Pulvermaar Tephra. (A) Under magnifier: dark grains are clinopyroxenes; colorless to pale greenish grains are olivines. (B) Rows: (1) Cpx = euhedral clinopyroxene (single mineral or agglomerate); O = olivine. Columns: (1) and (2) orthogonal positions in plane polarized light to highlight pleochroism; (3), (4), (5) maximum illumination between cross polarized light to highlight birefringence with intercalated gypsum in (4) and (5).



**Figure 5:** Vertical distribution of vmm of the RT in Upper Pleistocene sequences in Belgium. *Legend:* 1 = Humiferous Complex of Remicourt (HCR); 2 = palaeosol; 3 = frequency of vmm of the RT; 4 = unconformity. Horizontal scale = quantity of vmm for varying weights of sediment from one site to another (see the original publications).

been highlighted (Fig. 5). The part above the main peak comes from sedimentary reworking and the lower part from bioturbation (Juvigné, 1977b).

In the loess region of Middle Belgium (Rocourt, Kesselt, Romont, Veldwezelt, Remicourt, Tongrinne – Fig. 5), the tephra is systematically associated with a humiferous pedocomplex known as the Humiferous Complex of Remicourt (HCR; Haesaerts et al., 1997, 2016). In High Belgium, two sites yielded detailed data for the RT stratigraphic distribution. In the Walou Cave, the distribution peak is located on top of a humiferous horizon (unit CV-1), overlying a complex sequence including palaeosols (units DI-BT, CV-3 and CV-2; Pirson and Juvigné, 2011). In the Wanlin brickyard, the peak of the RT was observed at the boundary between two palaeosols separated by a stony layer (Juvigné, 1979a).

Hence, it seems that the fallout of the RT has occurred during a period of (relatively strong) soil formation, with sufficiently intense biological activity to disseminate the RT into the underlying units. In all the cases, the peak is situated above the Rocourt Pedocomplex or its equivalent. In the most complete loessic sequences, the vmm concentration peak is generally in the Humiferous Complex of Remicourt (see Juvigné et al., 2013; Haesaerts et al., 2016). There are a few exceptions, however. In one of the sequences studied at Rocourt, the RT peak was found above the HCR, but in other sequences, it was observed inside the HCR. In two sequences from Kesselt, the peak is below the HCR, but on other sequences from the same site, the highest concentration of RT-vmm was found inside the HCR. At Tongrinne, the peak is located at the contact between the Rocourt Pedocomplex and the overlying HCR.

Regarding the age of the RT, according to Haesaerts et al. (2016), the Rocourt Pedocomplex is attributed to the Eemian interglacial and to the main part of the Weichselian Early Glacial (GS 25 to the lower half of GI 21, sensu Rasmussen et al., 2014). Therefore, again according

to Haesaerts et al. (2016), the overlying Humiferous Complex of Remicourt bearing the RT is attributed to the end of GI 21 (ca. 78–80 ka; Juvigné et al., 2013). It is worth mentioning here that following Antoine et al. (2016), the equivalent of HCR is correlated with GI 20 and GI 19. Based on this viewpoint, the age of the RT would be slightly younger, 76.5–70 ka, following the Rasmussen et al. (2014) chronology. We therefore suggest here to consider the range 80–70 ka for the age of the fallout in order to widen the scope for determining the ages of volcanoes or tephra in the Eifel Volcanic Field.

## **7. About Correlations with the Rocourt Tephra**

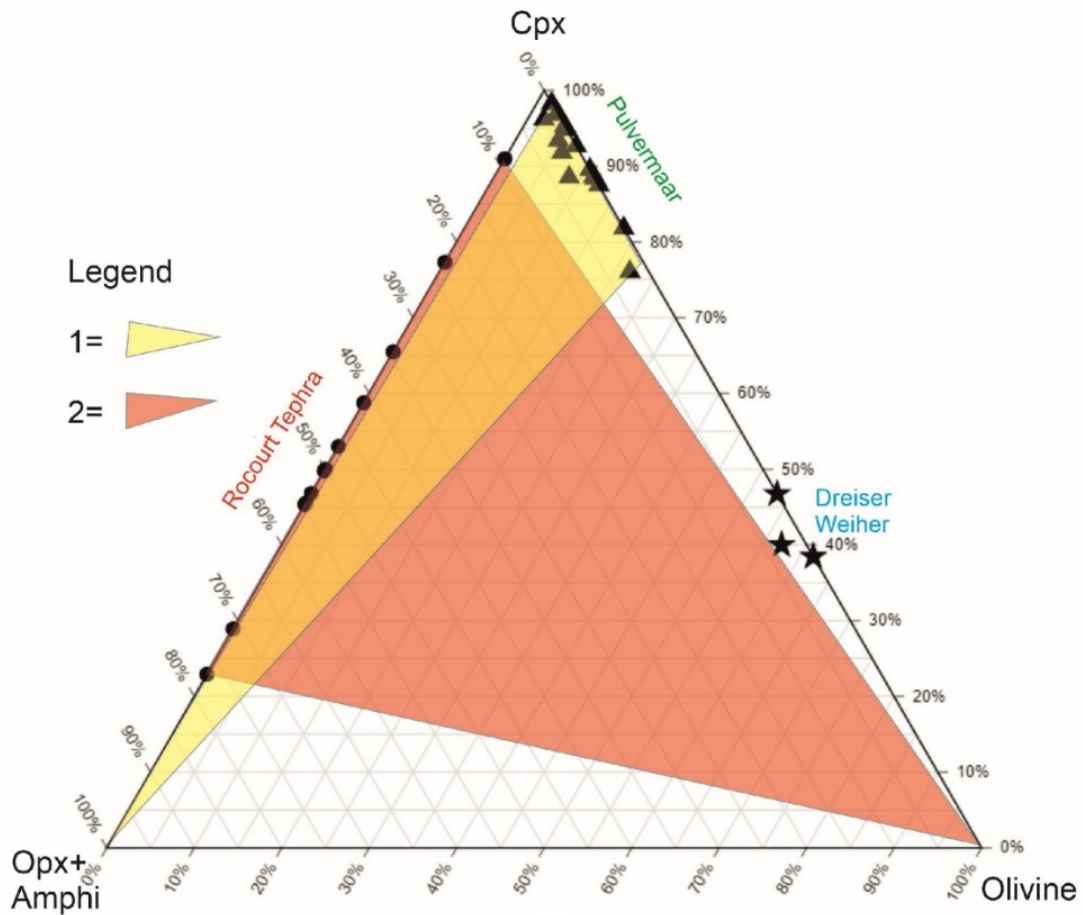
The identification of RT in some 35 sites in Belgium and in the neighboring regions of the Netherlands and the Lower Bay Rhine was based solely on determinations of vmm made under the microscope, and more particularly on the presence of enstatite as a marker mineral. The correlation of RT with products of the Dreiser Weiher (Gullentops and von der Hocht, 1998) or of the Pulvermaar (Förster et al., 2020) was rejected mainly by the geochemical fingerprints of pyroxenes (Juvigné et al., 2024). The detailed optical mineralogy data added in this paper allows us to argue in the same way. The photographs show fragments of megacrysts (without glass coatings) in the RT and euhedral clinopyroxenes with glass coating at Pulvermaar. Moreover, in a ternary diagram, there is no overlapping of the fields of the mineralogical associations of the three tephra (Fig. 6). It is also difficult to accept that weathering could have caused the disappearance of the amphiboles and enstatites from the Pulvermaar T. and the Dreiser Weiher T. and/or the olivines from the RT. Even if their total alteration were accepted, one could not obtain overlapping of the respective fields.

## **8. About the Presence of the RT at Schwalbenberg**

Fischer et al. (2021) report the presumed presence of the RT in the Schwalbenberg loess section (Middle Rhine valley). It is a centimeter-thick, coarse-grained tephra resting on a paleosol which could be the equivalent of the HCR in Belgium (see above). Unfortunately, the authors do not report any mineralogical or geochemical data concerning the RT in the Schwalbenberg loess. Their hypothesis calls for further investigations, because: (1) the site is on the road to the Inden and Garzweiler mines (Lower Bay Rhine) where Gullentops and von der Hocht (1998) found the RT; (2) the source volcano could be in the nearby East Eifel Volcanic Field, an unexpected hypothesis.

## **9. Conclusion**

Thus far, the volcano that has erupted the RT has not yet been identified. However, both the Pulvermaar and the Dreiser Weiher volcanoes can be ruled out, even though they have been suggested (erroneously) as sources in the literature. To propose the correlation of tephra with the RT, it is essential to refer not only to the chemical composition of the pyroxenes, but also to the association of volcanic mafic minerals which is characterised by high frequencies of



**Figure 6:** Comparisons of vmm associations of RT (black circles), Pulvermaar (black triangles) and Dreiser Weiher (black stars). *Legend:* 1 = possible distribution of the vmm association assuming the disappearance by alteration of brown amphiboles and enstatites in the proximal tephra of the Pulvermaar T.; 2 = possible distribution of the vmm association of the RT assuming the disappearance by alteration of olivines in all types of host sediments.

megacrystal fragments of enstatite and brown amphibole in the absence of olivine. The earlier age estimate for the RT of 78 to 80 ka (Juvigné et al., 2024)) is dependent on the stratigraphy of the loess compared to the INTIMATE curve, because the original stratigraphical position of the tephra is unknown. However, another stratigraphic-based suggestion for the age of RT is from 76.5 to 70 ka. Therefore, evidence of the presence of RT should not be sought only in the narrow age range of 78 to 80 ka but in the range 80-70 ka.

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## Further Information

### Author contributions

Etienne Juvigné: conceptualization, formal analysis, methodology, project administration, original draft writing, writing revision; André Pouclet, Jacques-Marie Bardintzeff, Stéphane Pirson: formal analysis, methodology, original draft writing, writing revision.

### Conflicts of interest

The authors declare that there is no conflict of interest.

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